A wide, flat, gravelly beach stretches across the foreground and middle ground, meeting a calm body of water on the left. The background features a dense line of green trees and some buildings under a clear blue sky with a few wispy clouds. The overall scene is bright and clear, suggesting a sunny day.

# EXPLAINING EXTREME EVENTS OF 2018

## From a Climate Perspective

Special Supplement to the  
*Bulletin of the American Meteorological Society*  
Vol. 101, No. 1, January 2020

# EXPLAINING EXTREME EVENTS OF 2018 FROM A CLIMATE PERSPECTIVE

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COVER CREDIT: iStock.com/Alena Kravchenko—River Thames receded during a heatwave in summer 2018 in London, United Kingdom.

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THE EXCEPTIONAL IBERIAN HEATWAVE  
OF SUMMER 2018

D. BARRIOPEDRO, P. M. SOUSA, R. M. TRIGO, R. GARCÍA-HERRERA, AND A. M. RAMOS

August 2018 saw the warmest Iberian heatwave since that of 2003. Recent climate change has exacerbated this event making it at least  $\gg 1^{\circ}\text{C}$  warmer than similar events since 1950–83.

The summer of 2018 was exceptionally warm in Europe, with outstanding temperatures over widespread non-contiguous areas, including Scandinavia, central Europe, Iberia, and the British Isles (e.g., WMO 2019). Different from other extraordinary summers, extreme temperatures did not occur during the same weeks everywhere, hitting the British Isles in June, Scandinavia and central Europe in July, and southwestern Europe in August. Together, they yielded the warmest European summer of the last 519 years, above the record-breaking summers of 2003 and 2010, albeit by a small margin, as inferred from instrumental and proxy data (Fig. 1a). Although northern and central Europe captured the attention of the media, Spain and Portugal experienced the warmest August after that of 2003

(AEMET 2019; IPMA 2019). Temperature anomalies were more pronounced during daytime over southwestern Iberia, and Portugal saw its warmest month in maximum temperature (TX) since 1931. Heat peaked during 1–7 August 2018, when an exceptional heatwave caused four (two) out of the five warmest days of the twenty-first century in Portugal (Spain), with country-mean daily TX reaching  $41.6^{\circ}\text{C}$  ( $36.4^{\circ}\text{C}$ ). We use observational and reanalysis data for 1950–2018 to quantify recent changes in the intensity of this kind of events.

**METHODS.** We describe the exceptionality (Fig. 1) and changing risk (Fig. 2) of the 2018 Iberian heatwave by using daily TX from E-OBS at  $0.25^{\circ} \times 0.25^{\circ}$  for 1950–2018 (Cornes et al. 2018) and historical series from the European Climate Assessment and Dataset (ECA&D) (Klein Tank et al. 2002) and the Instituto Português do Mar e da Atmosfera (IPMA). The atmospheric circulation is described with daily geopotential height at 500 hPa (Z500) and  $2.5^{\circ} \times 2.5^{\circ}$  from the NCEP–NCAR reanalysis (Kalnay et al. 1996). We use the analog method, which infers the probability distribution of a target field from the atmospheric circulation during the event (Stott et al. 2016, and references therein). Flow analog days are defined from their root-mean-square differences (RMSD) with the actual Z500 anomaly field over  $20^{\circ}\text{W}$ – $10^{\circ}\text{E}$ ,  $32.5^{\circ}$ – $50^{\circ}\text{N}$ . We reconstructed the Iberian ( $10^{\circ}\text{W}$ – $3.5^{\circ}\text{E}$ ,  $36^{\circ}$ – $43.5^{\circ}\text{N}$ ) mean TX by randomly picking one of the 20 best analogs for each heatwave day (1–7 August). This process was repeated 5,000 times with circulation analogs of the present (1984–2017) and past (1950–83) subperiods separately, building flow-conditioned distributions of Iberian TX for two different “worlds.” Their comparison provides the overall changes in heatwave intensity, including those due to non-anthropogenic factors [see Sánchez-Benítez et al. (2018) for details].

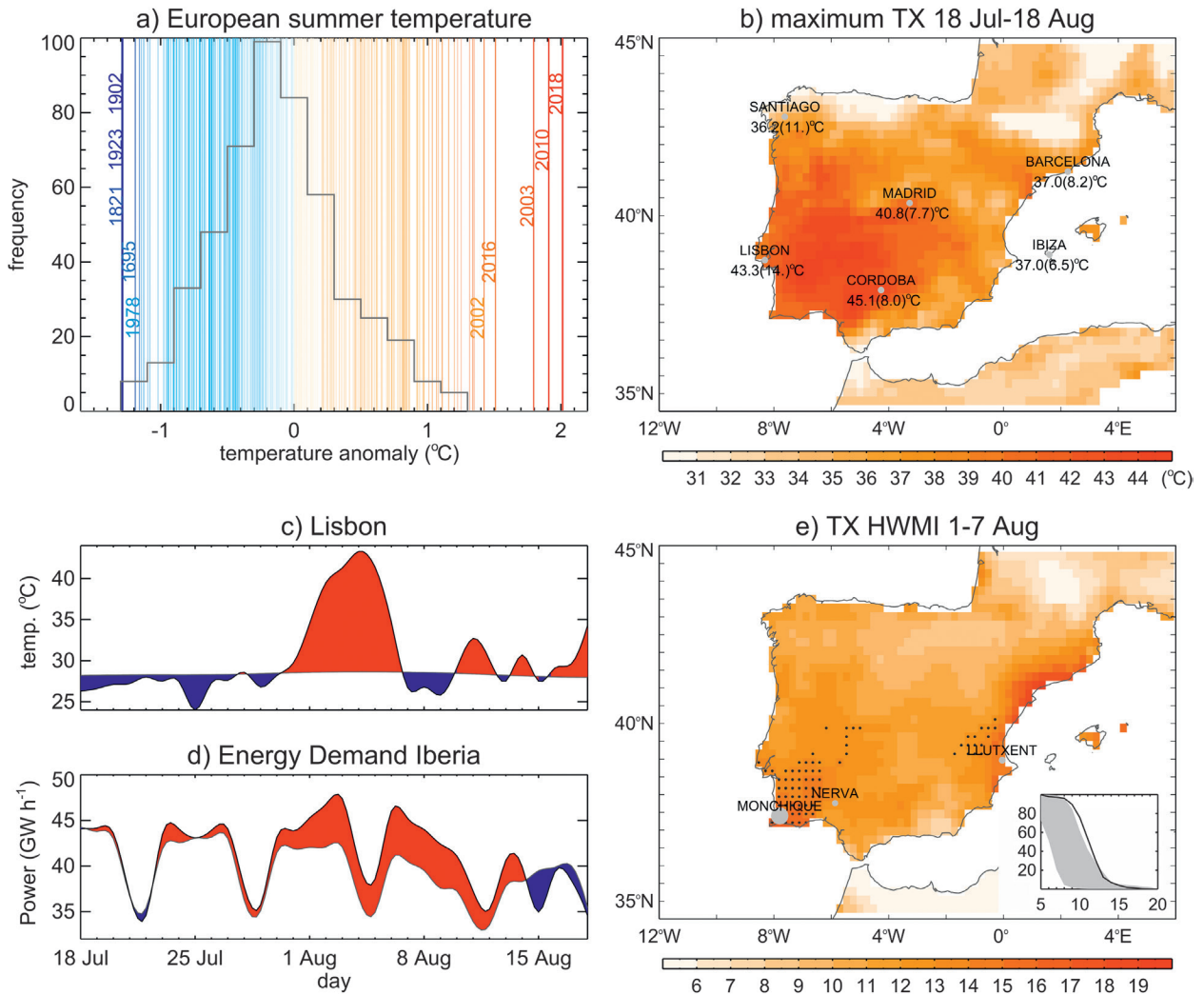
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**FIG. 1.** (a) European summer land temperature anomaly ( $^{\circ}\text{C}$ ; wrt 1981–2010) over  $25^{\circ}\text{W}$ – $40^{\circ}\text{E}$ ,  $35^{\circ}$ – $70^{\circ}\text{N}$  for 1500–2018 (lines) and its 1500–2000 frequency distribution (bars) using GISS (Hansen et al. 2010) for 1901–2018 and a multi-proxy reconstruction (Luterbacher et al. 2004). (b) Warmest daily TX for 18 Jul–18 Aug 2018 (shading) and selected stations (dots), with anomalies in parentheses ( $^{\circ}\text{C}$ ; wrt 1981–2010). Daily series of (c) Lisbon TX ( $^{\circ}\text{C}$ ) and (d) maximum electricity demand for Iberia ( $\text{GW h}^{-1}$ ), with red (blue) denoting periods above (below) average (1981–2010 and week-equivalent days of 2013–17, respectively). (e) HWMI (dimensionless) for 1–7 Aug 2018 and percentage of Iberia exceeding a given value (inset plot, with the 1981–2000 5th–95th percentile range in shading). Black dots indicate record-breaking values (wrt all 7-day intervals of 18 Jul–18 Aug) and gray dots indicate major fires.

**RESULTS.** Figure 1b shows the highest TX of the 18 July–18 August 2018 period, which is close to the warmest 31-day interval of the year over Iberia. TX climbed to  $46.8^{\circ}$  and  $46.6^{\circ}\text{C}$  in weather stations of Portugal and Spain (both on 4 August), close to their national records. Although the highest TX occurred in southern and western Iberia ( $\gg 40\%$  of the Portuguese stations broke their all-time records), unprecedented temperatures were also reported in central Iberia (e.g.,  $40.8^{\circ}\text{C}$ , Madrid), the Mediterranean coast (e.g.,  $39.8^{\circ}\text{C}$ , near Barcelona), and the Balearic Islands (e.g.,  $37.0^{\circ}\text{C}$ , Ibiza). Likewise, minimum temperatures

were exceptionally high, with more than 25% of the Portuguese stations setting absolute records and some Spanish locations reporting the warmest nights of the last century (e.g.,  $25.9^{\circ}\text{C}$ , Madrid). Tropical nights affected 50% of Portugal and extended to the Mediterranean coast (e.g.,  $>25^{\circ}\text{C}$ , Barcelona) during seven consecutive days.

The first week of August saw the warmest anomalies (Fig. 2a, shading), as illustrated by the time series of Lisbon (Fig. 1c), where TX surpassed  $40^{\circ}\text{C}$  for three days, breaking its previous record twice by a large margin ( $\gg 2^{\circ}\text{C}$  of exceedance). The atmospheric



of population in major touristic destinations, where the heat persisted the most. According to the Heat-wave Magnitude Intensity (HWMI) index [Russo et al. 2015; also see the online supplemental material (SM)], more than half of Iberia experienced extreme HWMI values (unprecedented in the southern half of Portugal and some Mediterranean areas; Fig. 1e), resulting in the most intense Iberian heatwave on 7-day time scales since 1950 after the 2003 episode (Table ES1).

Figure 2b shows the distribution of Iberian TX averaged for the heatwave period, as inferred from flow analogs of the past (blue boxplot) and present (red) climate. Present-day analogs explain almost 60% of the observed Iberian TX, the remaining being attributed to non-dynamical processes (e.g., feedbacks) and limited sampling. The comparison reveals that similar atmospheric conditions trigger warmer Iberian TX ( $\gg 1^\circ\text{C}$ ) now than in the recent past (i.e., the observed circulation would have caused a less severe heatwave in the past). This agrees with a warming and poleward trend of 2018-like Saharan intrusions, as reconstructed from flow analogs (see Fig. ES1 in the SM). Figure 2d quantifies how recent trends have changed the intensity of these Iberian heatwaves, by counting the fraction of replicated analogs with 7-day mean Iberian TX above a certain threshold in each subperiod. The flow-conditioned probability of experiencing Iberian heatwaves with TX anomalies above  $\gg 2.5^\circ\text{C}$  has doubled in just 35 years, equivalent to a fraction of attributable risk (FAR) of  $\gg 0.5$  (see the SM). Under the atmospheric circulation conditions of the 2018 heatwave, the chances of exceeding  $3^\circ\text{C}$  have risen by more than five times (FAR of 0.8).

**CONCLUSIONS AND DISCUSSION.** As the atmospheric circulation is constrained, the reported FAR should be attributed to thermodynamical changes (warming trend). However, flow analogs of the 2018 event show significant differences between the two subperiods, displaying smaller RMSD in the present than in the past (gray boxplots, Fig. 2b). Figure 2c (black line) confirms a significant ( $p < 0.05$ ) upward trend in the 1950–2018 frequency series of “good” flow analogs, defined as those days with RMSD below the 5th percentile of the event distribution. This trend may reflect dynamical (e.g., Z500 gradients) changes or thermodynamical effects (e.g., thermal Z500 rise). To address this, we repeated the analysis by removing the regional monthly mean trends of Z500 and TX. The resulting thermodynamically adjusted (TA) distributions for the two subperiods become much closer and the trend in the number of

good flow analogs is no longer significant at  $p < 0.05$  (gray line, Fig. 2c). Their difference has been added to the past distribution to estimate the contribution of dynamical changes (green line, Fig. 2d). Dynamical changes cannot explain the changing risk of Iberian TX anomalies. Therefore, regional warming is largely responsible for the FAR, particularly in the higher TX thresholds. Further studies are encouraged to pin down the key drivers and their contributing roles to the reported changes (e.g., land–atmosphere feedbacks).

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